CARNEGIE INSTITUTION Department of Terrestrial Magnetism

VT Earthquakes

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Introduction to VTs - Volcano-Seismic Event Classification

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Introduction to VTs - Timing of Occurrence

Generic Volcanic Earthquake Swarm Model

BCEAN





Introduction to VTs - Source process



How to generate a VT:

- 1. Increase shear stress
- 2. Decrease normal stress (directly)
- 3. Decrease normal stress (by increasing pore pressure)

VT Analysis Summary of Analyses

- Identify changes in space and/or time in:
 - Rates
 - Multiplets
 - Locations
 - Magnitudes
 - Focal Mechanisms
 - Shear-wave splitting
- Model or invert (based on Coulomb stress transfer)
 VT locations +- focal mechanisms +- split s-waves

Detect them (single-station)Count them (single-station)Automatically



From Guralp Systems

Detect them (single-station)Count them (single-station) - Manually



Case Study - Telica Volcano, Nicaragua



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← 2011 Eruption →



From Geirsson et al., in prep; Rodgers et al., in prep

- Look for waveform similarity (single station) multiplets
 - Identified through waveform cross-correlation
 - Similar waveforms implies similar location/source process
 - Reverse multiplets also sometimes observed



From Thelen et al. 2011

• Look for waveform similarity (single station) - multiplets

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From Mike West

Case Study - Bezymianny Volcano, Russia



BEEAN

- Locate them (4+ stations, velocity model)
- 1. Absolute location (Geiger 1910, 1912):
- Iterative least-squares approach
- Answer depends on spatial distribution of observations
- Requires a model of seismic velocities (P and S)



Locate them (4+ stations, velocity model)
2. Relative location (Double-Difference Method):





From Roman and Gardine 2013

Case Study - Long Valley, California



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From Shelly and Hill 2011

Case Study - Long Valley, California



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From Shelly and Hill 2011

Case Study - Soufriere Hills, Montserrat



- Most VTs locate within 1 km of vent
- Two distal VT clusters:
 - 1. St. George's Cluster Jul 1995 – Mar 1996
 - 2. NE Cluster August 5-6, 1995



From Roman et al. 2008

- VT Analysis - What can you do with a VT?
 - Determine the magnitude
 - Based on amplitudes, e.g., Richter magnitude

$$XMAG = \log(\frac{A}{2}) + [-B_1 + B_2 \log X^2] + G \qquad X = \sqrt{D^2 + Z^2}$$

Based on event duration,
 e.g., Coda magnitude (Lee and others, 1972; Lahr and others, 1975; Bakun and Lindh, 1977)

$$FMAG = C_1 + C_2 \log_{10} (F *_C) + C_3 D + C_4 Z + C_5 (\log_{10} (F *_C))^2$$

Case Study: VT Magnitudes - Kilauea Volcano, Hawai'i



From Wauthier et al., in prep

Case Study: VT Magnitudes - Kilauea Volcano, Hawai'i





From Wauthier et al., in prep

Case Study: VT Magnitudes - Kilauea Volcano, Hawai'i



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Determine/model the focal mechanism

First-motion polarities of rays leaving a hypocenter:



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Case Study: Focal Mechanisms - Mt. Spurr, Alaska



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From Roman et al. (2004)

Case Study: Focal Mechanisms - Mt Martin (Katmai), Alaska



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From O'Brien et al. (2012)

• Compare waveform on different seismometer components e.g., shear wave splitting, particle motions





Case Study - Soufriere Hills, Montserrat





From Roman et al. 2011

Models for the Origin of VTs - Numerical Models



Two induced stress regimes

 Compressive in walls of dike (hypocenters random in space)

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 Tension above propagating dike (hypocenters migrating ahead of dike tip)

(After Rubin and Pollard 1988, Ukawa and Tsukahara 1996)



(Roman and Cashman, 2006)

Models for the Origin of VTs - Observational Models



Volcano	Eruption	Study	~90° Rotation	Hypocenter Migration
Unzen, Japan	1990-1995	Umakoshi et al. (2001)	\checkmark	*
Usu, Japan	2000	Fukuyama et al. (2001)	\checkmark	*
Mt St. Helens, USA	1980-1986 2004	Barker and Malone (1991) Lehto et al. (2010)	\checkmark	*
Guagua Pichincha, Ecuador	1998	Legrand et al. (2002)	\checkmark	*
Crater Peak, Alaska	1992	Roman et al. (2004)	\checkmark	*
Soufriere Hills, Montserrat	1995-2007	Roman et al. (2006, 2008)	\checkmark	*
Redoubt Volcano, Alaska	2009	Gardine et al. (in prep)	\checkmark	*
Mt Etna, Italy	Multiple	e.g., Patane et al. (2003)	\checkmark	*
Teishi Knoll, Japan	1989	Ukawa and Tsukahara (20	01) 🔸	\checkmark
Miyake-jima, Japan	2000	Fukuyama et al. (2001)	*	\checkmark
Izu-Oshima, Japan	1987	Aramaki (1988)	*	\checkmark

Influence of Tectonic Stress - Patterns of VT seismicity

- In compressional environments, strongly deviatoric regional stresses can override volcanic stresses
- · Results in shadow zones where faults are locked



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Coulomb stress change (bars)



Influence of Tectonic Stress - Locations and FPS of VT earthquakes

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Intermediate Regional Compression Strong Regional Compression

Influence of Tectonic Stress - Locations and FPS of VT earthquakes

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- Results in shadow zones where faults are locked



-10.0 -9.0 -8.0 -7.0 -6.0 -5.0 -4.0 -3.0 -2.0 -1.0 .0 1.0 2.0 3.0 4.0 5.0 6.0 7.0 8.0 9.0 10.0

Influence of Tectonic Stress - Locations and FPS of VT earthquakes

• In isotropic environments, weak volcanic stresses dominate sense of fault slip



Influence of Tectonic Stress - Patterns of VT seismicity

- In compressional environments, strongly deviatoric regional stresses can override volcanic stresses
- · Results in shadow zones where faults are locked





- VTs are ubiquitous at active volcanoes and reflect the modification of stress in the crust by magma, fluid, and gas.
- Changes in properties of VTs through time can be indicative of changes in the magmatic system, and may form the basis for eruption forecasts.
- Patterns of VT seismicity are highly susceptible to the background state of stress in the crust, which can strongly influence the rates, locations, and focal mechanisms of VTs.